#### **Supplementary information**

### 1. Calculation of single-crystal strain rates

Single-crystal strain rates are calculated based on published data for San Carlos olivine crystal rheology (Table S2), including corrections for water content and oxygen fugacity. The general flow law is

$$\dot{\varepsilon} = A \sigma^n f O_2^m exp\left(\frac{E^* + PV^*}{RT}\right)$$
 Eq. S1

where  $\sigma$  is the differential stress,  $fO_2$  the oxygen fugacity, P the hydrostatic pressure, and T the temperature. R is the gas constant. A, n, m,  $E^*$ , and  $V^*$  are constants adjusted to previous experimental data (Table S2).

For a dry single crystal of olivine with the [110]<sub>c</sub> direction parallel to compression (promoting [100](010) dislocation slip), the crystal strain rate can be calculated according to

$$\dot{\varepsilon}_{[110]c} = B \left[ \dot{\varepsilon}_1 + ((\dot{\varepsilon}_2)^{-1} + (\dot{\varepsilon}_3)^{-1})^{-1} \right]$$
 Eq. S2

with  $\dot{\varepsilon}_1$ ,  $\dot{\varepsilon}_2$ ,  $\dot{\varepsilon}_3$  calculated using Eq. S1 and the parameters of Table S2. For a dry single-crystal, B = 1. For a wet single-crystal B = 4.13 (see below).

For a single crystal with the [011]<sub>c</sub> direction parallel to compression (promoting [001](010) dislocation slip), the corresponding strain rate is

$$\dot{\varepsilon}_{[011]c} = B \left[ \dot{\varepsilon}_1 + \dot{\varepsilon}_2 \right].$$
 Eq. S3

with  $\dot{\varepsilon}_1$  and  $\dot{\varepsilon}_2$  calculated using Eq. S1 and the parameters of Table S2. For a dry single-crystal, B = 1. For a wet single-crystal B = 24.7 (see below).

Finally, for a dry single crystal with the [101]<sub>c</sub> direction parallel to compression (promoting [100](001) and [001](100) slips together), the corresponding strain rate is

 $\dot{\varepsilon}_{[101]c} = B[((\dot{\varepsilon}_1)^{-1} + (\dot{\varepsilon}_2)^{-1})^{-1}]$  Eq. S4.

with  $\dot{\varepsilon}_1$  and  $\dot{\varepsilon}_2$  calculated using Eq. S1 and the parameters of Table S2. For a dry single-crystal, B = 1. For a wet single-crystal B = 11.3 (see below).

## 2. Effect of water

Wet olivine single crystals are weaker than dry crystals, respectively by factors of ~1.5, 2.5 and 2.0 on the differential stress when deformed along  $[110]_c$ ,  $[011]_c$  and  $[101]_c$  directions [*Mackwell et al.*, 1985; *Girard et al.*, 2013]. When calculating  $\dot{\epsilon}_{IC}$  for wet aggregates, these weakening factors translate into factors of 4.13, 24.7 and 11.3 on crystal strain rates, accounting for olivine stress exponent n = 3.5 in classical power laws [*Bai et al.*, 1991].

### 3. Supplemental References

Peslier, A.H., Woodland, A.B., Bell, D.R., Lazarov, M. (2010) Olivine water contents in the continental lithosphere and the longevity of cratons, *Nature* **467**, 78-81.

# **Supplementary Material**

**Figure S1**: plots of the 80-Ma and 20-Ma oceanic geotherms and the continental geotherm used in this study. Pressure increase with depth is calculated using the upper-mantle average density  $(3.35 \text{ g/cm}^3)$ , i.e. a 32.9 MPa/km vertical gradient.



**Figure S2**: same as Figure 3.b, with  $\sigma = 50$  MPa.



**Table S1:** Rheological data for olivine aggregates deformed in axial compression at high pressure and temperature.  $\dot{\epsilon}_{Agg}$  are raw experimental data.  $\dot{\epsilon}_{IC}$  are calculated from Eq. 3. For the unbuffered experiments, we assume either the iron-wüstite (IW) or the fayalite-magnetite-quartz (FMQ) buffer oxygen fugacity ( $fO_2$ ).

Run #	Р	$\Delta P$	Т	$\Delta T^{*}$	$\sigma$	$\Delta\sigma^*$	$fO_2$	$fO_2$	έ <sub>Agg</sub>	$\Delta \dot{\epsilon}_{Agg}$	$\log(\dot{\epsilon}_{Agg}/\dot{\epsilon}_{IC})$	$\log(\dot{\epsilon}_{Agg}/\dot{\epsilon}_{IC})$	$\Delta log(\dot{\epsilon}_{Agg}/\dot{\epsilon}_{IC})$
	GPa	GPa	Κ	Κ	MPa	MPa	atm	atm	$s^{-1}$	s <sup>-1</sup>			
Durham et al. [2009] : Dry. Ni-NiO buffered						Ni-NiO				Ni-NiO			
140	4.2	0.05	1540	25	250	50	6.13E-09	-	1.50E-05	5.80E-07	1.18	-	0.44
144(1)	2.7	0.05	1477	25	300	50	1.33E-09	-	1.90E-05	7.60E-07	1.39	-	0.42
144(2)	4.8	0.05	1455	25	500	50	7.52E-10	-	1.60E-05	5.20E-07	1.63	-	0.39
145(1)	4.9	0.05	1453	25	250	50	7.14E-10	-	1.90E-06	7.50E-08	1.83	-	0.47
145(2)	4.9	0.05	1451	25	340	50	6.77E-10	-	3.20E-05	1.30E-06	2.61	-	0.43
Hilairet et al. [2012]: Wet unbuffered					IW	FMQ			IW	FMQ			
D0912-1	4.2	0.2	1373	25	1385	220	4.86E-14	2.90E-10	6.70E-06	1.50E-07	0.69	-0.16	0.48
D1040-2	7.9	0.2	1375	25	2895	400	5.20E-14	3.08E-10	1.10E-05	5.00E-08	0.51	0.19	0.49
D1040-3	6.2	0.2	1375	25	2910	500	5.20E-14	3.08E-10	1.49E-05	2.00E-07	0.36	-0.19	0.50
D1062-1	5.0	0.2	1506	25	76	25	2.85E-12	1.19E-08	3.44E-06	2.00E-08	2.92	2.42	0.61
D1062-3	5.8	0.2	1506	25	1465	260	2.85E-12	1.19E-08	1.24E-05	6.00E-07	-0.80	-1.25	0.45
D1040-1	6.8	0.2	1593	25	169	25	2.83E-11	9.69E-08	1.27E-05	9.00E-08	0.54	0.00	0.47
D0912-2	3.8	0.2	1600	25	212	50	3.36E-11	1.14E-07	3.42E-05	2.90E-06	0.29	-0.24	0.52
D0912-3	4.3	0.2	1600	25	215	60	3.36E-11	1.14E-07	1.40E-05	4.00E-07	1.47	1.01	0.40
Bollinger et al. [2014]: Wet unbuffered					IW	FMQ			IW	FMQ			
81	7.2	0.1	1373	25	382	212	4.86E-14	2.90E-10	3.80E-05	1.00E-06	4.06	3.66	0.95
81	6.0	0.1	1373	25	493	160	4.86E-14	2.90E-10	2.70E-05	1.00E-06	3.32	2.74	0.65
82	10.6	0.1	1373	25	831	285	4.86E-14	2.90E-10	2.60E-05	1.00E-06	3.17	2.99	0.70
87	5.9	0.2	1373	25	737	183	4.86E-14	2.90E-10	1.80E-05	6.00E-07	2.51	1.91	0.57
267	4.2	0.2	1373	25	800	217	4.86E-14	2.90E-10	4.30E-05	1.00E-06	2.34	1.49	0.58
267	4.9	0.2	1373	25	917	595	4.86E-14	2.90E-10	2.40E-05	7.00E-07	2.07	1.32	1.07
80	4.0	0.1	1473	25	262	143	1.11E-12	5.05E-09	3.40E-05	5.00E-06	2.24	1.64	0.90
88	8.6	0.2	1473	25	884	170	1.11E-12	5.05E-09	1.60E-05	6.00E-07	1.08	0.80	0.49
89	5.7	0.2	1473	25	535	172	1.11E-12	5.05E-09	1.30E-05	1.00E-06	1.21	0.74	0.61
89	5.9	0.2	1473	25	504	191	1.11E-12	5.05E-09	1.10E-05	2.00E-06	1.28	0.82	0.69
90	7.9	0.2	1473	25	796	190	1.11E-12	5.05E-09	1.70E-05	8.00E-07	1.16	0.85	0.53
81	4.2	0.1	1573	25	341	88	1.71E-11	6.11E-08	5.80E-05	2.00E-06	0.60	0.07	0.50
82	9.0	0.1	1573	25	163	83	1.71E-11	6.11E-08	2.50E-05	5.00E-06	2.52	2.18	0.85
87	5.2	0.2	1573	25	203	203	1.71E-11	6.11E-08	2.40E-05	3.00E-06	1.32	0.82	1.56
87	6.1	0.2	1573	25	276	113	1.71E-11	6.11E-08	2.50E-05	2.00E-06	1.13	0.66	0.71
88	7.8	0.2	1673	25	272	137	1.89E-10	5.49E-07	2.80E-05	3.00E-06	0.28	-0.22	0.82
89	5.5	0.1	1673	25	83	98	1.89E-10	5.49E-07	4.40E-05	6.00E-06	1.65	1.08	1.82
89	4.8	0.2	1673	25	248	137	1.89E-10	5.49E-07	5.30E-05	1.00E-06	-0.16	-0.74	0.88
90	7.2	0.1	1673	25	118	90	1.89E-10	5.49E-07	4.50E-05	1.00E-06	1.61	1.09	1.20

 $^{*}\Delta T$  is the estimated uncertainty on nominal T;  $\Delta \sigma$  provided by N. Hilairet for *Hilairet et al.* [2012] data, and arbitrarily set to 50 MPa for *Durham et al* [2009] data. Note that, within uncertainties, all data points but one verify  $\log(\hat{\epsilon}_{Agg}/\hat{\epsilon}_{IC}) \ge 0$ , i.e.  $\hat{\epsilon}_{Agg}/\hat{\epsilon}_{IC} \ge 1$  in good agreement with Eq. 3;  $\hat{\epsilon}_{Agg}/\hat{\epsilon}_{IC} < 1$  may occasionally result from the lower-bound approach used here to calculate the intracrystalline strain rate (overestimation of  $\hat{\epsilon}_{IC}$ ). See text for further explanations.

1	Table S2.	Flow-law	parameters <sup>†</sup>	for orthop	yroxene-b	uffer SC	olivine	single of	crystals	3
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From raw . $\varepsilon = A \delta \int O_2 exp\left(\frac{-RT}{RT}\right)$								
Compression direction		A (MPa <sup>-n</sup> atm <sup>-m</sup> s <sup>-1</sup> )	п	т	E* (kJ/mol)	V* (cm <sup>3</sup> /mol)		
	Ė1	0.02	3.5	0.36	230	12		
[110] <sub>c</sub> <sup>††</sup>	έ <sub>2</sub>	1.3 ×10 <sup>22</sup>	3.5	0.10	1000	12		
	Ė3	1.2	3.5	0.15	290	12		
[011]c <sup>††</sup>	Ė <sub>1</sub>	2.1 ×10 <sup>4</sup>	3.5	0.02	540	3		
	Ė2	5.2 ×10 <sup>5</sup>	3.5	0.23	540	3		
[101] c <sup>††</sup>	Ė2	0.65	3.5	0.33	250	10.7		
	έ <sub>2</sub>	5.3 ×10 <sup>11</sup>	3.5	0.06	690	10.7		

Flow law:  $\dot{\epsilon} = A \sigma^n f \Omega_c^m \rho rn \left( \frac{E^* + PV^*}{E^*} \right)$ 

<sup>†</sup>A, n, m and E\* from Bai et al. [1991], V\* from Raterron et al. [2009; 2012]; <sup>††</sup> $\dot{\epsilon}_{[110]c} = \dot{\epsilon}_1 + ((\dot{\epsilon}_2)^{-1} + (\dot{\epsilon}_3)^{-1})^{-1}; \dot{\epsilon}_{[011]c} = \dot{\epsilon}_1 + \dot{\epsilon}_2; \dot{\epsilon}_{[101]c} = ((\dot{\epsilon}_1)^{-1} + (\dot{\epsilon}_2)^{-1})^{-1};$ in wet conditions, strain rates must be multiplied by 4.13, 24.7 and 11.3, respectively 

5 [Mackwell and Kohlstedt, 1985; Girard et al., 2013].