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## Surface Geosciences (Pedology) Variation of the water-retention properties of soils:

# Validity of class-pedotransfer functions

Hassan Al Majou<sup>a</sup>, Ary Bruand<sup>a,\*</sup>, Odile Duval<sup>b</sup>, Isabelle Cousin<sup>b</sup>

<sup>a</sup> Institut des sciences de la Terre d'Orléans (ISTO), UMR 6113, CNRS, université d'Orléans, 1A, rue de la Férollerie, 45072 Orléans cedex 2, France

<sup>b</sup> Unité de science du sol, INRA, centre de recherche d'Orléans, BP 20619, 45166 Olivet cedex, France

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#### Abstract

Water-retention properties of soils vary according to soil characteristics, and the understanding of their variation remains controversial. Numerous pedotransfer functions (ptfs) that enable prediction of the water-retention properties of soils were developed, but their validity was poorly discussed. In this study, we compare the performances of textural and texturo-structural class ptfs with more sophisticated class and continuous ptfs developed using the same set of soils. We showed that the former led to prediction performances that are better than, or similar to, those recorded with the more sophisticated class and continuous ptfs that are quite easy to establish are potentially worthwhile tools for predicting the water-retention properties of soils, particularly at scales for which semi-quantitative or qualitative basic soil characteristics, such as the texture, are the only characteristics available. More generally, our results pointed out that the discussion of ptfs performance should refer to those recorded with easy to establish ptfs, thus enabling to quantify how much prediction bias and precision can be gained when increasing the complexity of ptfs and, consequently, the number and quality of predictors required. *To cite this article: H. Al Majou et al., C. R. Geoscience 339 (2007).* 

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## Résumé

Variation des propriétés de rétention en eau des sols : validité des classes de pédotransfert. Les propriétés de rétention en eau des sols varient en fonction de leur composition, et elles sont encore largement discutées. De nombreuses fonctions de pédotransfert (fpt) permettant de les prédire ont été développées, mais leur validité n'a été que rarement discutée. Dans cette étude, nous comparons les performances de classes de fpt texturales et texturo-structurales développées en utilisant un même jeu de données. Nous montrons que les classes de fpt conduisent à des performances de prédiction qui sont meilleures que, ou similaires à celles enregistrées avec les fpt plus sophistiquées étudiées par ailleurs dans cette étude. Ainsi, les classes de fpt texturales et texturo-structurales qu'il est aisé d'établir sont potentiellement des outils utiles pour la prédiction des propriétés de rétention en eau des sols, en particulier aux échelles auxquelles seules des données semi-quantitatives ou qualitatives, comme la texture, sont disponibles. Plus généralement, nos résultats mettent en évidence le fait que les performances des fpt devraient être discutées en prenant comme référence celles enregistrées avec des fpt faciles à établir, comme les classes de fpt texturales. En procédant ainsi, il est alors possible

\* Corresponding author.

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E-mail address: Ary.Bruand@univ-orleans.fr (A. Bruand).

d'apprécier le gain de performance en termes de biais et de précision quand on complexifie les fpt et que l'on accroît le nombre et la qualité des caractéristiques de sols requises. *Pour citer cet article : H. Al Majou et al., C. R. Geoscience 339 (2007).* © 2007 Académie des sciences. Published by Elsevier Masson SAS. All rights reserved.

Keywords: Texture; Bulk density; Horizon; Structure; Prediction bias; Prediction precision

Mots clés : Texture ; Densité apparente ; Horizon ; Structure ; Biais de prédiction ; Précision

## 1. Introduction

Understanding water-retention properties of soil remains a major issue in soil science. Because of the growing demand for soil hydraulic properties, a common solution has been to use pedotransfer functions (ptfs) that relate basic soil properties that are considered as easily accessible to the less often measured soil properties, such as hydraulic properties [1]. A huge number of ptfs was developed over the last three decades and we are facing today the continuous development of ptfs of increasing complexity, with very little or no information about the potential increase in the prediction quality. There is some information available about the performance of continuous ptfs [10,17], very little about the performance of class ptfs [14,17], and even less about the compared performance of these two types of ptfs [15]. The aim of this study is to show that the variation of water-retention properties can be predicted by using stratification based on information about particle-size distribution and structure. We show also that the quality of the prediction is similar to or better than that achieved with much more sophisticated ptfs, despite what is usually admitted.

## 2. Materials and methods

## 2.1. The ptfs developed in the literature

Most ptfs published in the literature are continuous pedotransfer functions (continuous ptfs), i.e. mathematical continuous functions between the water content at discrete values of potential or the parameters of a unique model of water-retention curve and the basic soil properties (mostly particle-size distribution, organic carbon content and bulk density) [12,17]. Besides these continuous ptfs that enable continuously the prediction of water content at particular water potentials [13] or estimation of the parameters of models of the water-retention curve [5,10,17], there are class pedotransfer functions (class ptfs) that received little attention, because their accuracy is considered as limited [15].

The existing class ptfs often provide average water contents at particular water potentials, or one average water-retention curve for every texture class [3,11]. Due to the range in particle-size distribution, clay mineralogy, organic matter content, and structural development within each texture class, water-retention properties for individual soils were considered as varying considerably [16]. Despite their possible inaccuracies, class ptfs enable the prediction based on successive stratification using soil characteristics. Moreover, class ptfs are easy to use because they require little soil information and are well adapted to the prediction of water-retention properties over large areas [9,15,16].

## 2.2. The soils studied

Class and continuous ptfs were developed using a set of 320 horizons, comprising 90 topsoils (from 0 to 30 cm depth) and 230 subsoil horizons (> 30 cm depth) collected in Cambisols, Luvisols, Planosols, Albeluvisols, Podzols, and Fluvisols [8] located mainly in the Paris basin and secondarily in the western coastal marshlands and Pyrenean piedmont plain. A set of 107 horizons comprising 39 topsoil and 68 subsoil horizons was constituted in order to test the ptfs established. These horizons were collected in Cambisols, Luvisols and Fluvisols [8] located in the South of the Paris basin. Basic characteristics and water-retention properties of the horizons were determined as described earlier by Bruand and Tessier [2] (Fig. 1, Table 1). Their bulk density  $(D_b)$  was measured by using cylinders 1000 cm<sup>3</sup> in volume when the soil was near to field capacity.

## 2.3. Analysis of the ptfs performance

In order to discuss the global validity of the ptfs, most studies used the root-mean-square error (*RMSE*), which is also called root-mean-square deviation or root-mean-square residual [17]. Because the *RMSE* varies according to both the prediction bias and precision, we computed the mean error of prediction (*MEP*) that enables discussion of the prediction bias alone, on the



Fig. 1. Triangle of texture used ( $\mathbf{a}$ ), texture of the horizons used to develop the class and continuous ptfs ( $\mathbf{b}$ ) and texture of those used to test their validity ( $\mathbf{c}$ ).

Fig. 1. Triangle de texture utilisé (a), texture des horizons utilisés pour développer les classes de fpt et les fpt continues (b) et texture des horizons utilisés pour discuter leur validité (c).

one hand, and the standard deviation of prediction (*SDP*) that enables discussion of the prediction precision alone, on the other hand. We computed *MEP* and *SDP* for the whole water potentials as follows:

$$MEP = \frac{1}{l' \cdot l} \sum_{j=1}^{l'} \sum_{i=1}^{l} (\theta_{p,j,i} - \theta_{m,j,i})$$
$$SDP = \left\{ \frac{1}{l' \cdot l} \sum_{j=1}^{l'} \sum_{i=1}^{l} [(\theta_{p,j,i} - \theta_{m,j,i}) - MEP]^2 \right\}^{1/2}$$

.,

where  $\theta_{p,j,i}$  is the predicted water content at potential *i* for the horizon *j*,  $\theta_{m,j,i}$  is the measured water content at potential *i* for the horizon *j*, and *l* is the number of water potentials for each horizon (l = 7 in this study) and *l'* is the number of horizons ( $l' \leq 107$  in this study). The *MEP* corresponds to the bias and indicates whether the ptfs overestimated (positive) or underestimated (negative) the water content, whereas *SDP* measures the precision of the prediction.

In order to discuss the validity of the ptfs at the different water potentials, we computed also the mean

Table 1 Characteristics of the horizons of the data set used to develop the ptfs and of the test data set

#### Tableau 1

Caractéristiques des horizons de l'ensemble de données utilisées pour développer les fpt et de celles utilisées pour en discuter la validité

	Particle	Particle size distribution (%)			CaCO <sub>3</sub>	CEC	D <sub>b</sub>	Volumetric water content (cm <sup>3</sup> cm <sup>-3</sup> )						
	$<2\mu m$	2–50 µm	50–2000 µm	$(g kg^{-1})$	$(g kg^{-1})$	$(\operatorname{cmol}_{c} \operatorname{kg}^{-1})$	(g cm <sup>-5</sup> )	$\theta_{1.0}$	$\theta_{1.5}$	$\theta_{2.0}$	$\theta_{2.5}$	$\theta_{3.0}$	$\theta_{3.5}$	$\theta_{4.2}$
Horizons	used to	establish cla	ass and continu	ious ptfs ( <i>i</i>	i = 320)									
mean	28.9	46.2	24.9	5.7	65	14.3	1.53	0.350	0.335	0.316	0.289	0.257	0.220	0.179
s.d.	15.1	20.8	23.9	4.9	189	8.0	0.15	0.067	0.065	0.070	0.070	0.075	0.074	0.070
min.	1.9	2.8	0.1	0.0	0.0	0.8	1.00	0.123	0.100	0.080	0.056	0.048	0.033	0.013
max.	92.9	82.1	90.1	28.8	982	52.8	1.84	0.606	0.596	0.586	0.558	0.510	0.462	0.370
Horizons	used to	test the ptfs	s (n = 107)											
mean	30.2	40.6	29.2	6.6	38	15.8	1.51	0.356	0.332	0.312	0.287	0.261	0.224	0.202
s.d.	15.4	24.3	28.6	5.3	134	10.8	0.13	0.075	0.079	0.082	0.084	0.086	0.083	0.080
min.	1.9	4.1	1.6	0.0	0.0	0.6	1.10	0.161	0.121	0.099	0.072	0.045	0.041	0.033
max.	78.7	80.3	91.8	28.2	656	50.2	1.77	0.534	0.498	0.482	0.457	0.440	0.396	0.369

error of prediction (MEP') and the standard deviation (SDP') of prediction at every water potential as follows:

$$MEP' = \frac{1}{l'} \sum_{j=1}^{l'} (\theta_{p,j} - \theta_{m,j})$$
$$SDP' = \left\{ \frac{1}{l'} \sum_{j=1}^{l'} [(\theta_{p,j} - \theta_{m,j}) - MEP']^2 \right\}^{1/2}$$

## 3. Results and discussion

## 3.1. The class and continuous ptfs developed

The class ptfs developed in this note were established according to the texture (textural class ptfs) in the CEC triangle [4] and then according to both that texture and  $D_{\rm b}$  (texturo-structural class ptfs). The resulting class ptfs corresponded to the average water content at seven water potentials, which was computed within every class of texture (textural class ptfs) (Table 2) and every class combining both texture and  $D_{\rm b}$  (texturo-structural

Table 2 Textural class ptfs developed Tableau 2 Classes de fpt texturales développées class ptfs) (Table 3). More complex class ptfs were established by fitting the van Genuchten's model [6] on the arithmetic mean value of  $\theta$  at the different values of water potential using the RETC code [7] for every class of texture (VG texture class ptfs) according to the CEC triangle [4] and to the type of horizon (topsoil and subsoil) (Table 4).

Continuous ptfs were also developed. They correspond to multiple regression equations as follows:

$$\theta = a + (b \times \text{\%Cl}) + (c \times \text{\%Si}) + (d \times \text{\%OC}) + (e \times D_b)$$

with  $\theta$  the volumetric water content at a given water content, *a*, *b*, *c*, *d* and *e* the regression coefficients, %CI and %Si, respectively, the clay and silt contents, %OC, the organic carbon content, and  $D_b$ , the bulk density (Table 5). Other continuous ptfs were developed as earlier done by Wösten et al. [16] for the parameters of the van Genuchten's model, using multiple regression equations (VG continuous ptfs, Table 6). For every horizon, the parameters of the van Genuchten's model were computed using the RETC code [7].

	Volumetric water content (cm <sup>3</sup> cm <sup>-3</sup> )											
	$\overline{ heta_{1.0}}$	$\theta_{1.5}$	$\theta_{2.0}$	$\theta_{2.5}$	$\theta_{3.0}$	$\theta_{3.5}$	$\theta_{4.2}$					
Very fine $(n = 15)$	0.455	0.437	0.424	0.402	0.385	0.357	0.322					
Fine $(n = 60)$	0.399	0.388	0.373	0.351	0.331	0.301	0.254					
Medium fine $(n = 96)$	0.356	0.342	0.327	0.298	0.254	0.210	0.173					
Medium $(n = 117)$	0.334	0.320	0.302	0.273	0.242	0.203	0.156					
Coarse $(n = 32)$	0.249	0.224	0.181	0.149	0.120	0.100	0.076					

Table 3 Texturo-structural class ptfs developed Tableau 3

Clas	sses	de	fpt	texturo-structural	es	déve	loppées
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		Volumetric water content $(cm^3 cm^{-3})$										
		$\theta_{1.0}$	$\theta_{1.5}$	$\theta_{2.0}$	$\theta_{2.5}$	$\theta_{3.0}$	$\theta_{3.5}$	$\theta_{4.2}$				
Very Fine $(n = 15)$	$1.10 \le D_{\rm b} < 1.30$	0.498	0.473	0.451	0.423	0.405	0.371	0.330				
	$1.30 \le D_{\rm b} < 1.50$	0.459	0.439	0.428	0.405	0.385	0.352	0.328				
	$1.50 \le D_{\rm b} < 1.70$	0.359	0.359	0.361	0.353	0.347	0.340	0.294				
Fine $(n = 60)$	$1.00 \le D_{\rm b} < 1.20$	0.519	0.499	0.494	0.461	0.431	0.373	0.281				
	$1.20 \le D_{\rm b} < 1.40$	0.452	0.443	0.421	0.385	0.373	0.340	0.271				
	$1.40 \le D_{\rm b} < 1.60$	0.391	0.378	0.361	0.344	0.321	0.289	0.250				
	$1.60 \le D_{\rm b} < 1.80$	0.338	0.334	0.325	0.307	0.291	0.275	0.244				
Medium Fine $(n = 96)$	$1.20 \le D_{\rm b} < 1.40$	0.348	0.338	0.323	0.291	0.232	0.188	0.153				
	$1.40 \le D_{\rm b} < 1.60$	0.359	0.343	0.328	0.298	0.258	0.211	0.175				
	$1.60 \le D_{\rm b} < 1.80$	0.353	0.345	0.329	0.303	0.263	0.230	0.190				
Medium ( $n = 117$ )	$1.20 \le D_{\rm b} < 1.40$	0.354	0.337	0.314	0.278	0.245	0.193	0.140				
	$1.40 \le D_{\rm b} < 1.60$	0.346	0.329	0.310	0.275	0.235	0.193	0.146				
	$1.60 \le D_{\rm b} < 1.80$	0.320	0.307	0.293	0.270	0.248	0.214	0.167				
	$1.80 \le D_{\rm b} < 2.00$	0.296	0.289	0.274	0.266	0.258	0.231	0.186				
Coarse $(n = 32)$	$1.40 \le D_{\rm b} < 1.60$	0.241	0.210	0.164	0.135	0.106	0.093	0.075				
	$1.60 \le D_{\rm b} < 1.80$	0.253	0.231	0.188	0.156	0.126	0.103	0.077				

## 3.2. Validity of the class ptfs

The textural class ptfs underestimated very slightly the water retained ( $MEP = -0.003 \text{ cm}^3 \text{ cm}^{-3}$ ) when they were applied to the test dataset without any other stratification than according to the texture. There was no decrease in the prediction bias with the texturo-

#### Table 4

Parameters of the van Genuchten's model corresponding to the VG textural class ptfs developed according to the type of horizon (topsoil and subsoil)

Tableau 4

Paramètres du modèle de van Genuchten correspondant aux classes de ptf VG texturales développées en fonction du type d'horizon (horizon de surface et horizon de subsurface)

	$\theta_{\rm r}$	$\theta_{\rm s}$	α	n	т
Topsoils					
Coarse	0.025	0.397	1.0592	1.1530	0.1327
Medium	0.010	0.428	0.4467	1.1000	0.0909
Medium fine	0.010	0.465	0.6860	1.1027	0.0931
Fine	0.010	0.477	0.6153	1.0652	0.0612
Very Fine	0.010	0.587	5.9433	1.0658	0.0617
Subsoils					
Coarse	0.025	0.367	1.0535	1.1878	0.1581
Medium	0.010	0.388	0.1851	1.0992	0.0903
Medium fine	0.010	0.416	0.1611	1.0978	0.0891
Fine	0.010	0.437	0.1334	1.0632	0.0594
Very Fine	0.010	0.472	0.0745	1.0499	0.0475

structural class ptfs ( $MEP = -0.004 \text{ cm}^3 \text{ cm}^{-3}$ ), but the bias was already very small with the textural class ptfs studied. However, the precision was slightly better with the texturo-structural class ptfs (SDP =  $0.043 \text{ cm}^3 \text{ cm}^{-3}$ ) than with the textural class ptfs  $(SDP = 0.045 \text{ cm}^3 \text{ cm}^{-3})$  (Fig. 2a and b). Compared to the textural class ptfs, the VG textural class ptfs showed similar performance. The bias was very small  $(MEP = 0.002 \text{ cm}^3 \text{ cm}^{-3})$  and the precision poor  $(SDP = 0.045 \text{ cm}^3 \text{ cm}^{-3})$ , as recorded for the textural class ptfs (Fig. 2c). The comparison of the class ptfs performance at every value of water potential showed small bias  $(-0.008 \le MEP' \le 0.007 \text{ cm}^3 \text{ cm}^{-3})$  except for  $\theta_{4,2}$  for the textural and texturo-structural class ptfs  $(MEP' = -0.020 \text{ and } -0.019 \text{ cm}^3 \text{ cm}^{-3})$  and for  $\theta_{1,0}$  for the VG class ptfs ( $MEP' = 0.014 \text{ cm}^3 \text{ cm}^{-3}$ ), for which it was greater (Table 7). This comparison showed also poor precision for the three class ptfs studied, whatever the water potential  $(0.040 < SDP' < 0.047 \text{ cm}^3 \text{ cm}^{-3})$ .

## 3.3. Validity of the continuous ptfs

When applied to the test data set, the continuous ptfs lead to very small bias ( $MEP = -0.003 \text{ cm}^3 \text{ cm}^{-3}$ ) and showed poor precision ( $SDP = 0.039 \text{ cm}^3 \text{ cm}^{-3}$ ). Results showed a greater bias with the VG continuous ptfs ( $MEP = -0.008 \text{ cm}^3 \text{ cm}^{-3}$ ) and similar poor precision ( $SDP = 0.039 \text{ cm}^3 \text{ cm}^{-3}$ ) than with the Table 5

Regression coefficients and coefficient of determination  $R^2$  recorded for the continuous ptfs developed Tableau 5

coefficients de regression et coefficients de determination re-	Coefficients	s de r	régression	et c	oefficients	de	détermination	$R^2$	enregistrés	pour	les	ptf	continues	déve	lop	pées
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	Water potential	Water potential (hPa)													
	-10	-33	-100	-330	-1000	-3300	-15000								
a	0.4701***	0.3556***	0.2620***	0.1301***	0.0184	-0.0504	-0.0786**								
b	$0.0026^{***}$	$0.0029^{***}$	0.0034***	0.0038***	$0.0045^{***}$	$0.0047^{***}$	$0.0045^{***}$								
с	$0.0006^{***}$	$0.0008^{***}$	$0.0012^{***}$	$0.0012^{***}$	$0.0008^{***}$	$0.0005^{***}$	$0.0003^{***}$								
d	-0.0006	-0.0002	0.0002	0.0010	$0.0017^{***}$	$0.0012^{**}$	0.0004								
е	-0.1447***	-0.0939***	$-0.0647^{***}$	-0.0084	$0.0398^{*}$	$0.0697^{***}$	$0.0710^{***}$								
$R^2$	0.59	0.64	0.69	0.74	0.77	0.82	0.86								

 $\theta = a + (b \times \% \text{Cl}) + (c \times \% \text{Si}) + (d \times \% \text{OC}) + (e \times D_b)$  with  $\theta$  volumetric water content at a given water content. \*\*\* P = 0.001. \*\* P = 0.01.

Table 6

VG continuous ptfs developed for the parameters of the van Genuchten's model

Tableau 6

Relations correspondant aux fpt VG continues développées pour les paramètres du modèle de van Genuchten

$$\begin{split} \theta_{\rm s} &= 1.1658 - 0.0032 \times C - 0.4737 \times D + 2 \times 10^{-7} \times S^2 - 0.0001 \times OC^2 + 0.0373 \times C^{-1} + 0.0131 \times S^{-1} - 0.0072 \times \ln(S) + 0.00003 \\ &\times OC \times C + 0.0022 \times D \times C - 0.0002 \times D \times OC - 0.0001 \times S \ (R^2 = 0.95) \\ \alpha^* &= 25.61 + 0.0439 \times C + 0.1129 \times S + 1.1914 \times OC + 32.21 \times D - 10.48 \times D^2 - 0.0009 \times C^2 - 0.0146 \times OC^2 \\ &- 0.3781 \times OC^{-1} - 0.0178 \times \ln(S) - 0.1032 \times \ln(OC) - 0.1 \times D \times S - 0.6001 \times D \times OC \ (R^2 = 0.26) \\ n^* &= -15.29 - 0.0659 \times C + 0.0115 \times S - 0.2115 \times OC + 12.33 \times D - 1.3578 \times D^2 + 0.0006 \times C^2 + 0.0031 \times OC^2 + 4.0005 \times D^{-1} \end{split}$$

 $+ 2.2003 \times S^{-1} + 0.1643 \times OC^{-1} - 0.1205 \times \ln(S) + 0.2693 \times \ln(OC) - 9.9367 \times \ln(D) + 0.003 \times D \times C + 0.0694 \times D \times OC \ (R^2 = 0.35) \times OC^{-1} + 0.1643 \times OC^{-1} - 0.1205 \times \ln(S) + 0.2693 \times \ln(OC) - 9.9367 \times \ln(D) + 0.003 \times D \times C + 0.0694 \times D \times OC \ (R^2 = 0.35) \times OC^{-1} + 0.1643 \times OC^{-1} + 0.1$ 

 $\theta_s$  is a model parameter,  $\alpha^*$ ,  $n^*$  are transformed model parameters in the Mualem–van Genuchten equations; C = percentage clay (i.e., percentage < 2 µm); S = percentage silt (i.e., percentage between 2 µm and 50 µm); OC = organic carbon g kg<sup>-1</sup>; D = bulk density.

continuous ptfs (Fig. 2d and e). The comparison of the continuous ptfs performance at every value of water potential showed small bias for the continuous ptfs  $(-0.006 \le MEP' \le 0.005 \text{ cm}^3 \text{ cm}^{-3})$ , except for  $\theta_{4.2}$  ( $MEP' = -0.022 \text{ cm}^3 \text{ cm}^{-3}$ ). For the VG continuous ptfs, the bias was greater for six water potentials with absolute value of  $MEP' \le 0.020 \text{ cm}^3 \text{ cm}^{-3}$ , except for  $\theta_{1.5}$  ( $MEP' = 0.004 \text{ cm}^3 \text{ cm}^{-3}$ ) (Table 7). The precision was poor for the simple and VG continuous ptfs

Table 7

Validity of the continuous and class ptfs according to the water potential

Tableau 7

Validité des classes de fpt et des fpt continues aux différentes valeurs de potentiel de l'eau

 $(0.030 \le SDP' \le 0.044 \text{ cm}^3 \text{ cm}^{-3})$ , but results showed that *SDP* decreased with the water potential.

#### 3.4. Comparison of the class- and continuous ptfs

Results showed very little difference between the ptfs studied. The bias recorded was small  $(-0.008 \le MEP \le 0.002 \text{ cm}^3 \text{ cm}^{-3})$  and the greatest absolute value of bias was recorded with the VG

	Volumet	ric wate	r content (	$(cm^3 cm^{-3})$	<sup>3</sup> )									
	Mean er	ror of p	rediction (	MEP')	Standard deviation of prediction (SDP')									
	$\theta_{1.0}$	$\theta_{1.5}$	$\theta_{2.0}$	$\theta_{2.5}$	$\theta_{3.0}$	$\theta_{3.5}$	$\theta_{4.2}$	$\theta_{1.0}$	$\theta_{1.5}$	$\theta_{2.0}$	$\theta_{2.5}$	$\theta_{3.0}$	$\theta_{3.5}$	$\theta_{4.2}$
Textural class ptfs	-0.006	0.004	0.003	0.001	-0.004	-0.001	-0.020	0.046	0.046	0.044	0.045	0.047	0.044	0.042
Texturo-structural class ptfs	-0.006	0.002	0.002	0.001	-0.005	-0.002	-0.019	0.042	0.042	0.041	0.043	0.045	0.044	0.041
VG class ptfs	0.014	0.007	-0.003	-0.008	-0.007	0.007	0.002	0.045	0.045	0.045	0.046	0.046	0.043	0.040
Continuous ptfs VG continuous ptfs	$-0.006 \\ 0.012$	0.001 0.004	$0.005 \\ -0.008$	$0.001 \\ -0.017$	$-0.003 \\ -0.020$	$0.002 \\ -0.008$	$-0.022 \\ -0.016$	0.044 0.044	0.044 0.041	0.040 0.038	0.039 0.039	0.036 0.035	0.032 0.033	0.030 0.032



Fig. 2. Validity of the textural class ptfs (**a**), texturo-structural class ptfs (**b**), VG textural class ptfs (**c**), continuous ptfs (**d**), and VG continuous ptfs (**e**) developed.

Fig. 2. Validité des classes de fpt texturales (a), texturo-structurales (b), et VG texturales (c), ainsi que des fpt continues (d) et VG continues (e).

continuous ptfs ( $MEP = -0.008 \text{ cm}^3 \text{ cm}^{-3}$ ). On the other hand, the precision was poor ( $0.039 \le SDP \le 0.045 \text{ cm}^3 \text{ cm}^{-3}$ ), the greatest precision being recorded with the two types of continuous ptfs studied. If the VG continuous ptfs led to the greatest precision ( $SDP = 0.039 \text{ cm}^3 \text{ cm}^{-3}$ ), they led also the greatest bias value ( $MEP = -0.008 \text{ cm}^3 \text{ cm}^{-3}$ ).

## 4. Conclusion

Our results showed that textural class ptfs led to prediction performance that are similar to those recorded with more sophisticated class ptfs and with continuous ptfs. Thus without knowing the particle-size distribution, organic carbon content and bulk density as required by most ptfs, we can predict the water-retention properties with similar prediction quality by using the texture alone. Our results showed also that use of both texture and bulk density slightly increases the precision when compared to the precision recorded with the textural class ptfs. Finally, we showed also that class ptfs, including very simple ptfs, should be still considered as useful tools for predicting the water-retention properties of soils, particularly at scales for which semi-quantitative or qualitative basic soil characteristic such as the texture are the only characteristics available. More generally, our results pointed out that discussion of ptfs performance should refer to those recorded with simple ptfs, thus enabling to quantify how much prediction bias and precision can be gained when increasing the complexity of ptfs and consequently the number and quality of predictors required.

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